

AMA2015: 19 January, 2015

Revision of Soil Thermodynamics and Its Impacts on Surface Meteorology in ORCHIDEE-LMDZ Coupled Model

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Outline

- ▶ **Background:** Soil thermal & hydrology in ORCHIDEE-11
- ▶ **Three revisions** in ORCHIDEE-11
 - I. Soil thermal property parameterization
 - II. Soil vertical discretization
 - III. Including soil heat convection by liquid water transfer
- ▶ **Conclusions**

Background:

Soil Thermal & Hydrology in ORCHIDEE-11

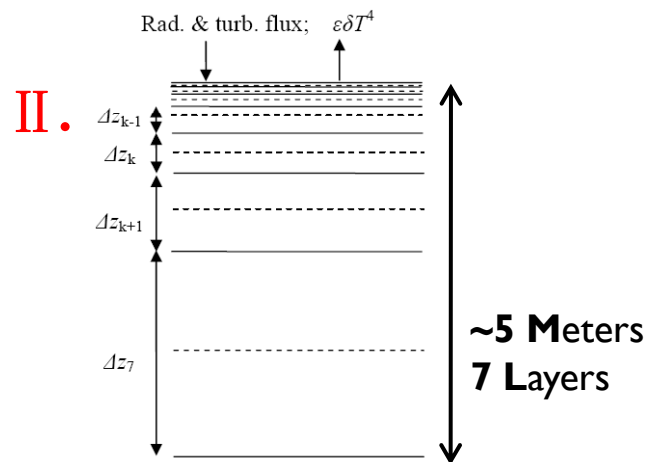
Currently: **I.** soil moisture changes with soil textures, but soil thermal properties do not; **II.** different soil vertical discretizations for moisture & temperature; **III.** soil heat convection process is neglected. **Energy imbalance !**

Soil Thermal [F. Hourdin, 1992]

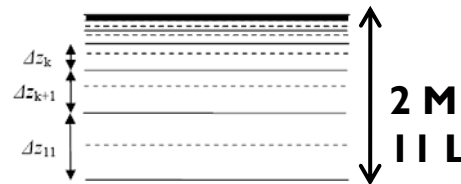
Soil Hydrology [P. de Rosnay, 1999]

Revisions

I. One soil texture.



Three soil textures:
Coarse, Medium, Fine



I. Soil Thermal

Properties change with soil textures & moisture.

II. Same soil vertical layers for temperature & moisture.

III. Coupling soil heat convection-conduction.

III.
$$C_p(\theta) \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\lambda(\theta) \frac{\partial T}{\partial z} \right]$$

C_p : soil heat capa. (J/m³/K),
 λ : thermal conduc. (W/m/K).

$$\frac{\partial \theta}{\partial t} = \frac{\partial q_L}{\partial z} - s$$

$$q_L = -D(\theta) \frac{\partial \theta}{\partial z} + K(\theta)$$

θ : soil moisture (m³/m³);
 q_L : liquid water flux (m/s);

I . Soil Thermal Properties in ORCHIDEE

Currently

Soil Heat Capacity

$$C_p(\theta) = C_{dry} + \frac{\theta - \theta_w}{\theta_f - \theta_w} \times (C_{wet} - C_{dry})$$

Soil Thermal Conductivity

$$\kappa(\theta) = \kappa_{dry} + \frac{\theta - \theta_w}{\theta_f - \theta_w} \times (\kappa_{wet} - \kappa_{dry})$$

θ : Volumetric soil moisture.

θ_f, θ_w : θ at field capacity & wilting point.

$C_{dry}, C_{wet}, \kappa_{dry}, \kappa_{wet}$ (prescribed).

(same thermal property for different soil textures).

Revised

$$C_p(\theta, st) = C_d(st) + \theta(st) \times C_w$$

θ : Vol. soil moist.; C_d : Dry capacity, Pielke [2002].

$$\kappa(\theta, st) = K_e(\theta, st) \times [\kappa_{sat}(st) - \kappa_{dry}(st)] + \kappa_{dry}(st)$$

$$\kappa_{dry}(st) = \frac{0.135 \times [1 - n_p(st)] \times 2700 + 64.7}{2700 - 0.947 \times [1 - n_p(st)] \times 2700}$$

$$\kappa_{sat}(st) = \left[(\kappa_q^{q(st)} \kappa_o^{1-q(st)}) \right]^{1-n_p(st)} \kappa_w^{n_p(st)}$$

$$K_e(\theta, st) = 0.7 \log S_r[\theta(st)] + 1.0$$

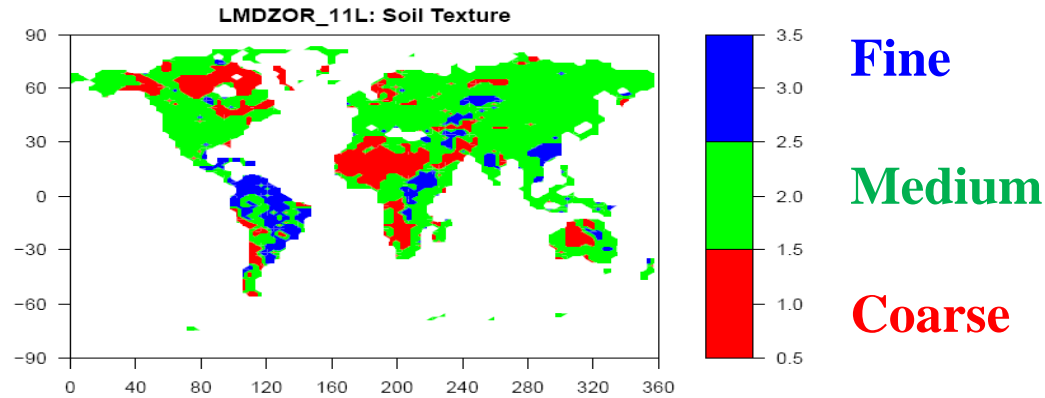
$\kappa_q, \kappa_o, \kappa_w$: κ for quartz, other minerals, water (prescribed).

n_p : porosity; S_r : Saturation degree; q : quartz.

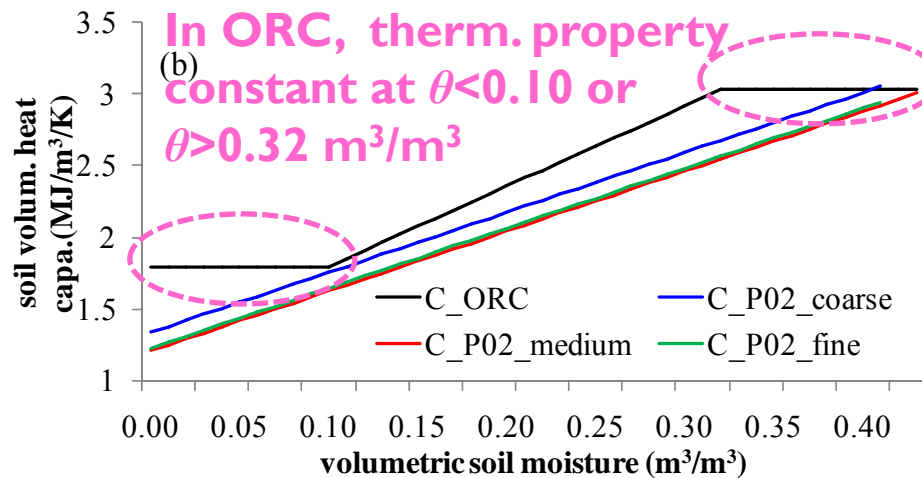
(thermal property varies with different soil textures)

Soil Thermal Properties with --- 3 Soil Texture Classes

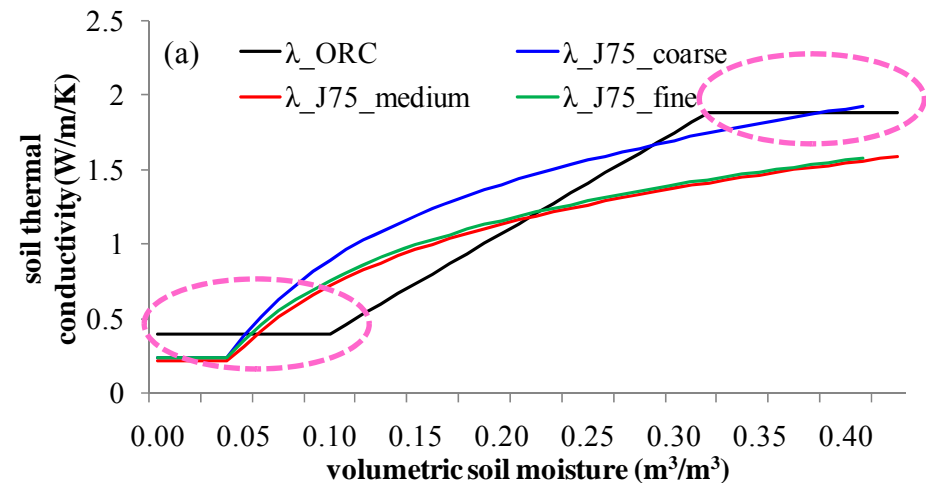
Three Soil Textures in ORCHIDEE-11



Soil Heat Capacity

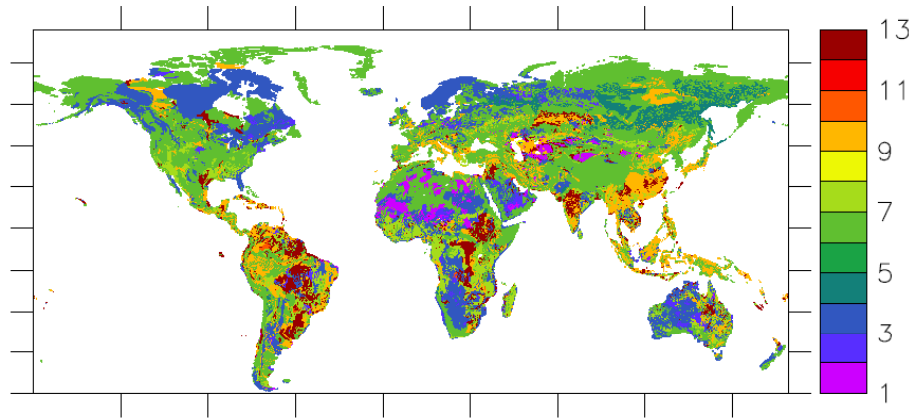


Soil Thermal Conductivity

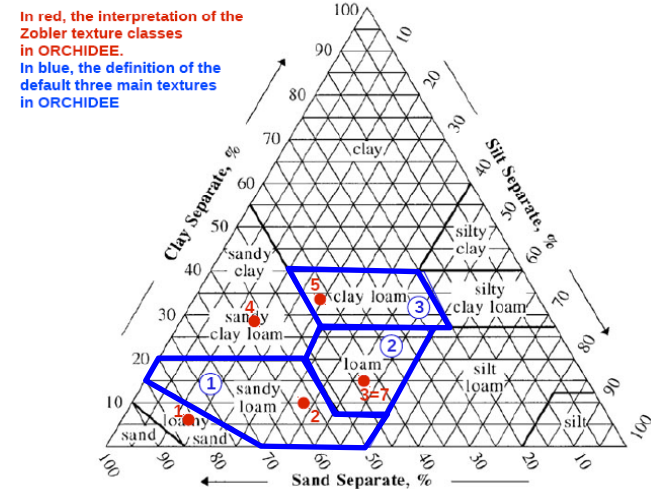


Soil Thermal Properties with --- USDA 12 Soil Texture Classes

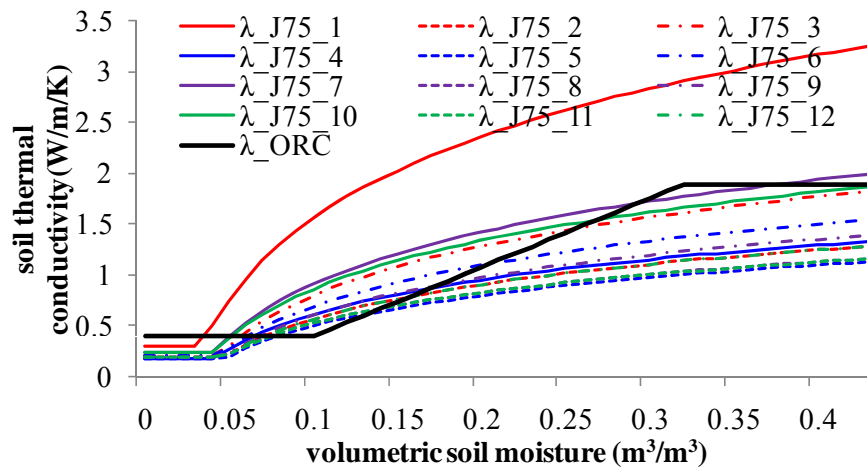
USDA 12 Soil Textures



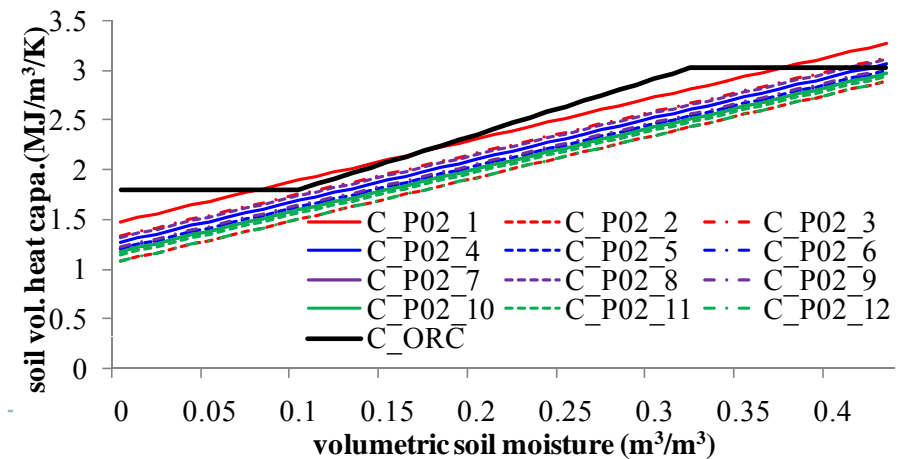
USDA Soil Texture Triangle



Soil Heat Capacity



Soil Thermal Conductivity



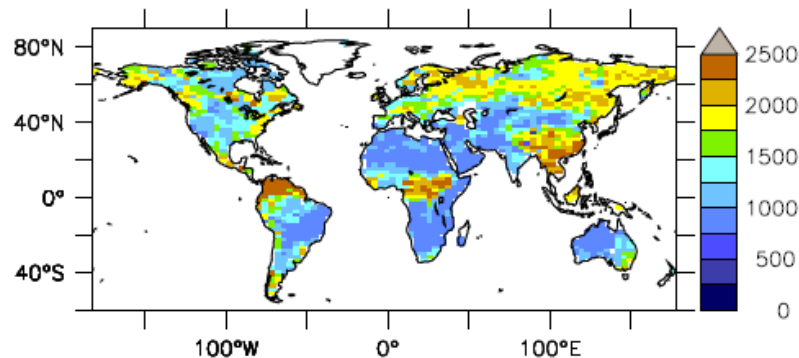
Effects of Soil Thermal Properties (3 textures) in ORC-LMDZ, JJA (3-Year AVE)

	Thermal property	Discretization	moisture	
CTL	Standard method	Standard	3 textures	Nudged
EXP	Revised (3 textures)			

Soil thermal inertia, $P = (\lambda C_p)^{0.5}$, ($W/m^2/K*s^{0.5}$), resistance of soil to temp. change.

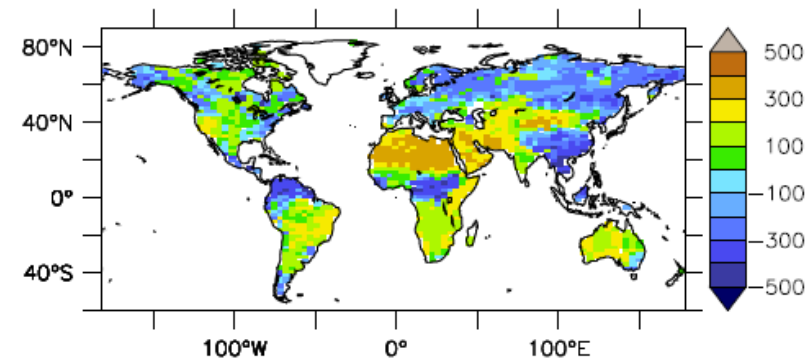
Standard

(i) SOILINER CTL($W/m^2/K*s^{0.5}$,JJA)

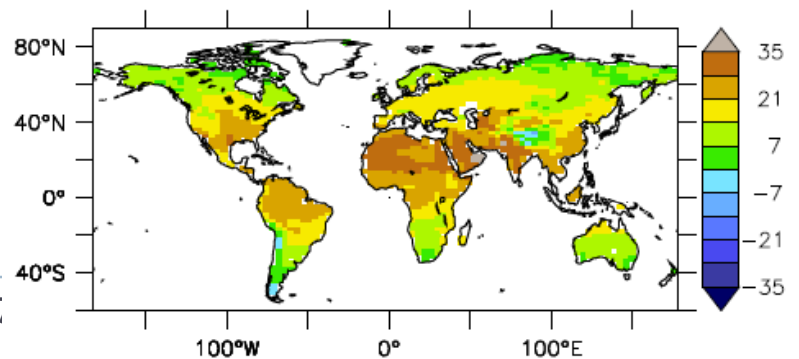


Revised - Standard

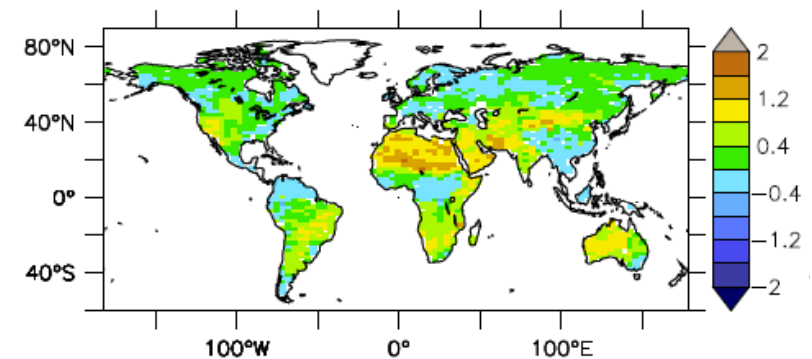
(j) SOILINER EXP-CTL($W/m^2/K*s^{0.5}$,JJA)



(k) TSURF CTL(Celsius,JJA)



(l) TSURF EXP-CTL(Celsius,JJA)

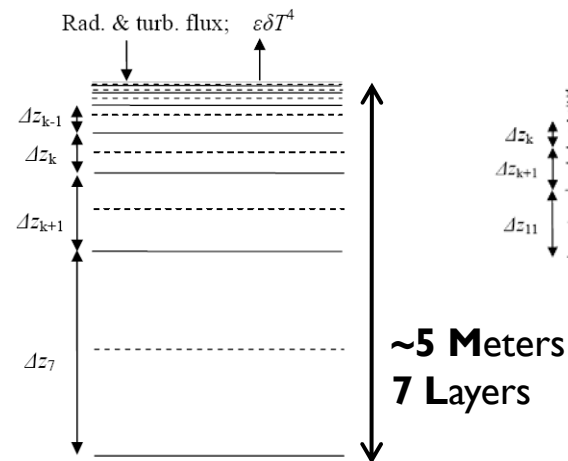


II. Soil Vertical Discretizations in ORCHIDEE

Currently: Different soil vertical discretizations for moisture & temperature; Energy imbalance !

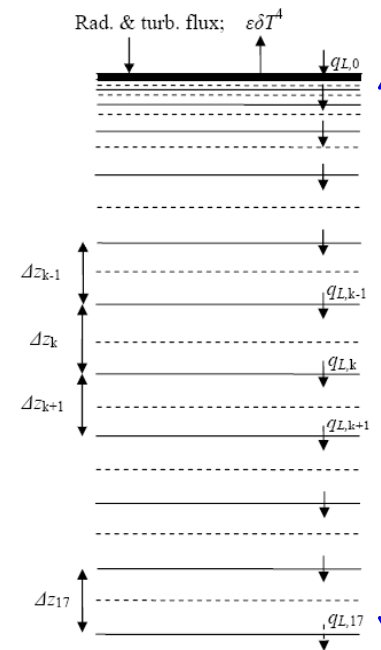
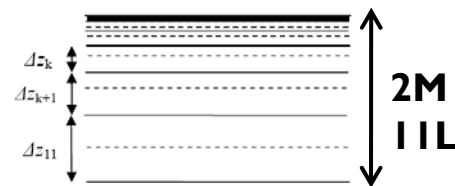
Revised: The same soil vertical discretization for moisture & temperature.

Soil Thermodynamics
[F. Hourdin, 1992]



BC: Zero heat flux.

Hydrology
[P. de Rosnay, 1999]



8M
17L

Above 2m:

same with standard hydro. model;

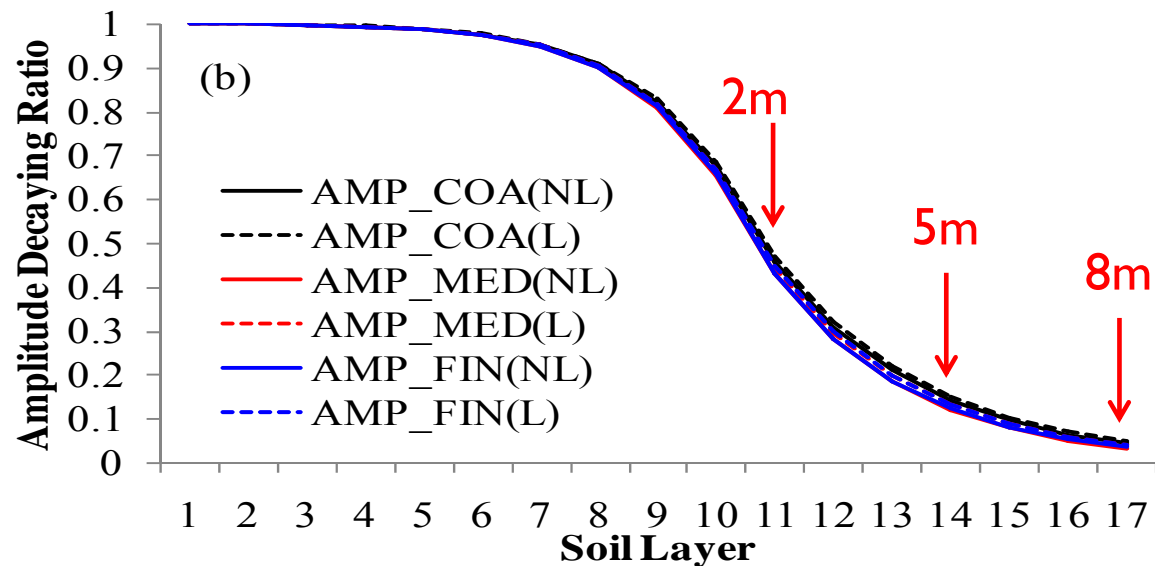
Below 2m:

increased by 1m for each layer.

Why Extend Soil Depth to 8M ?

- Diurnal/Annual cycle of temperature (solar/net radiation).
- The amplitude of soil temperature variation decreases with depth.
- To have enough soil depth to simulate the diurnal/annual cycles of soil temperature.

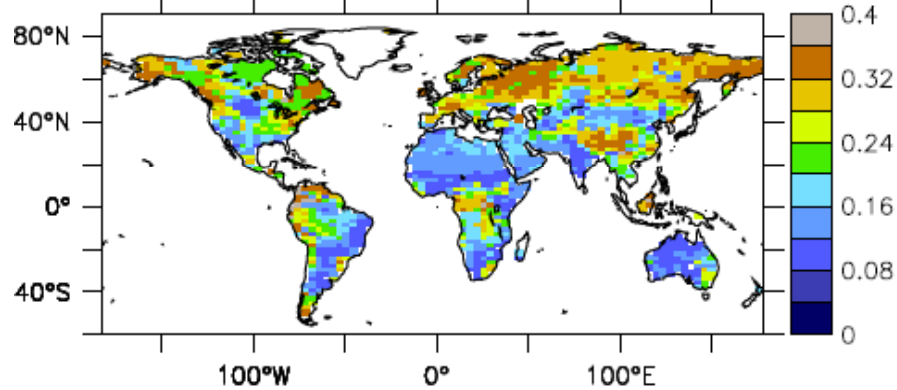
The decaying of temperature amplitude with soil depth →



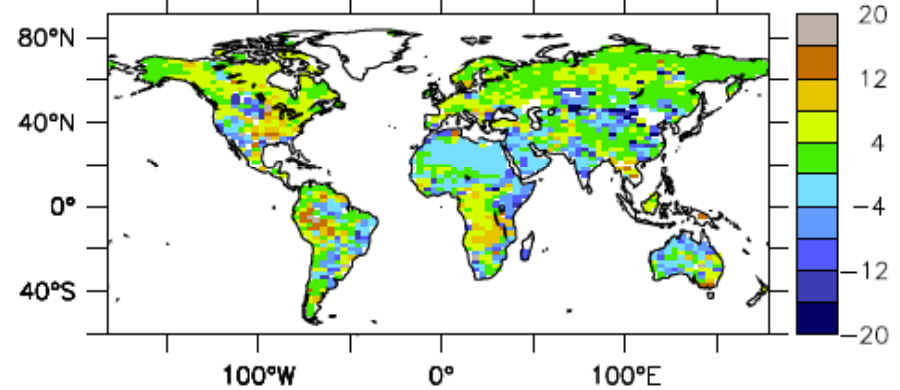
Soil Discretization Effects (3-Year)

Soil Discretization		Thermal Property	
CTL	Standard	Constant globally	Nudged
EXP	Revised (8 meters)		

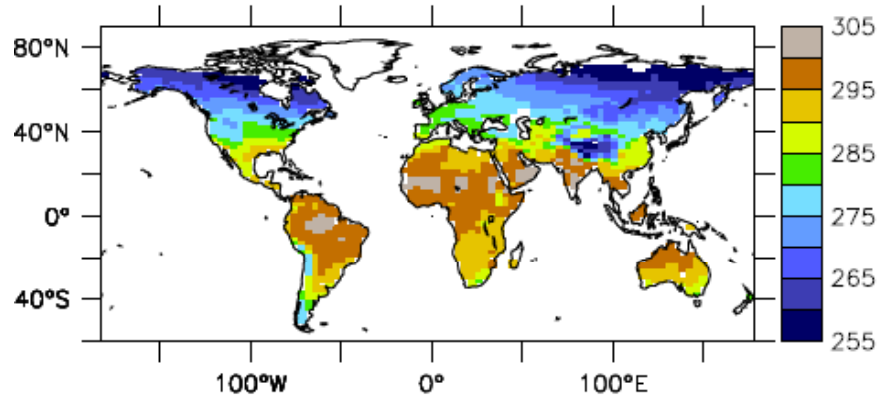
(a) VSMC 0–1.5M CTL(m³/m³,YR)



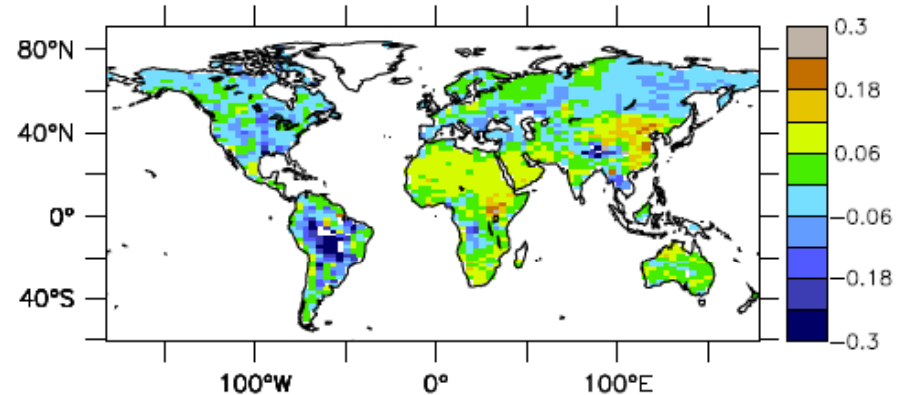
(b) VSMC 0–1.5M (EXP–CTL)/CTL(percent,YR)



(i) TAIR(K,YR)



(j) TAIR EXP–CTL(K, YR)



III. Soil Heat Conduction-Convection Coupling (Numerical)

Currently: soil heat convection is neglected, only conduction is considered.

Revised: Soil heat convection-conduction coupled.

$$C_p(\theta) \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\lambda(\theta) \frac{\partial T}{\partial z} \right]$$

C_p : soil heat capa. (J/m³/K),
 λ : thermal conduc. (W/m/K).

[F. Hourdin, 1992]

$$C_p(\theta, st) \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\lambda(\theta, st) \frac{\partial T}{\partial z} \right] - C_w \frac{\partial q_L T}{\partial z} - S$$

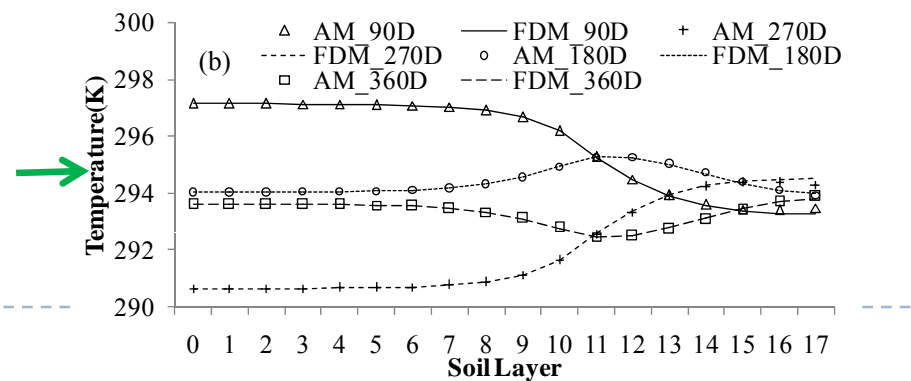
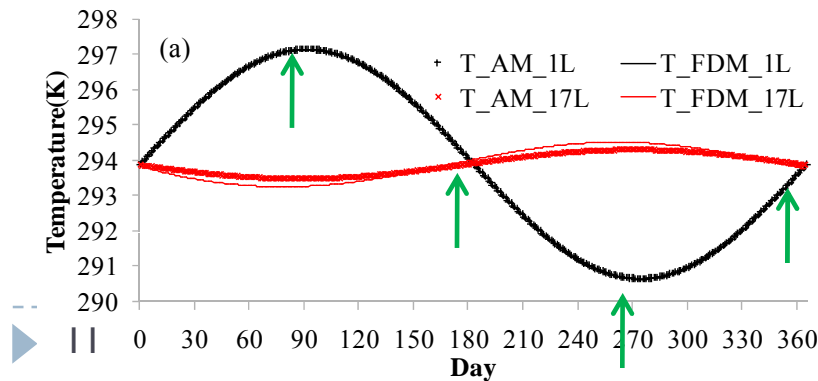
$$\frac{\partial \theta}{\partial t} = \frac{\partial q_L}{\partial z} - s$$

C_w : water heat capa. (J/m³/K),
 θ : soil moist. (m³/m³); q_L : liquid water flux (m/s)

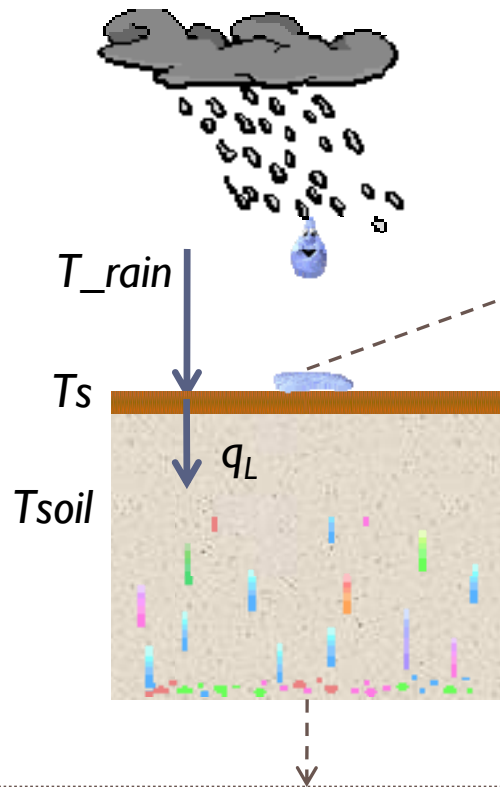
Evaluation of the revised model in 1-D case ($q_L=1E-7$ m/s, 8.6 mm/d) by analytical solutions.

Annual Cycle: Time series at different layers

Vertical profile at different time steps



Soil Heat Conduction-Convection Coupling (Boundary Conditions in ORCHIDEE-11)



Bottom BC: zero flux, $dT/dz=0$

Land Surface BC:

➤ Surface Energy Balance:

$$C_p \times dT_s/dt = R_{net} + LH + SH + G + \mathbf{H_{rain}}$$

$$\mathbf{H_{rain}} = C_w \times (T_{rain} - T_s) \times q_{L,0}$$

$$T_{rain} \approx T_{wetbulb} \text{ [Gosnell et al., 1995]}$$

$$T_{wetbulb} \sim f(T_{air}, Q_{air}, P_s)$$

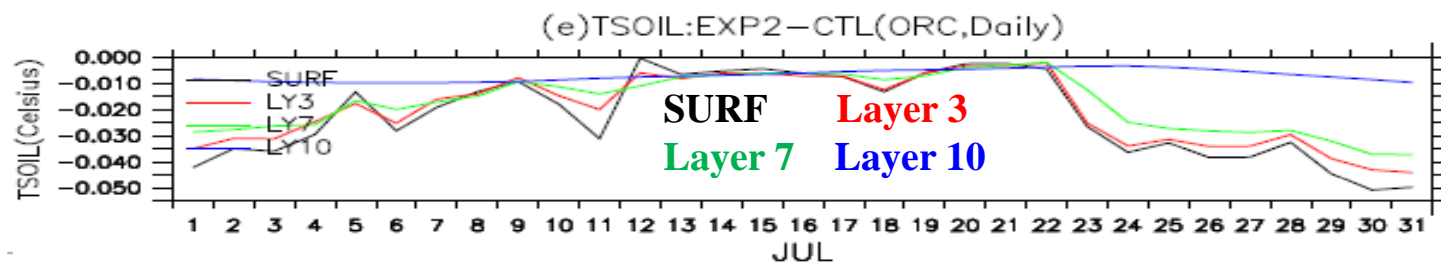
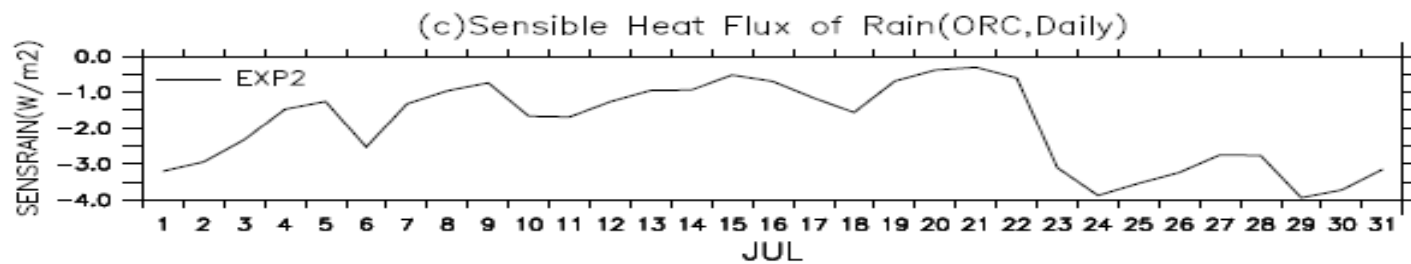
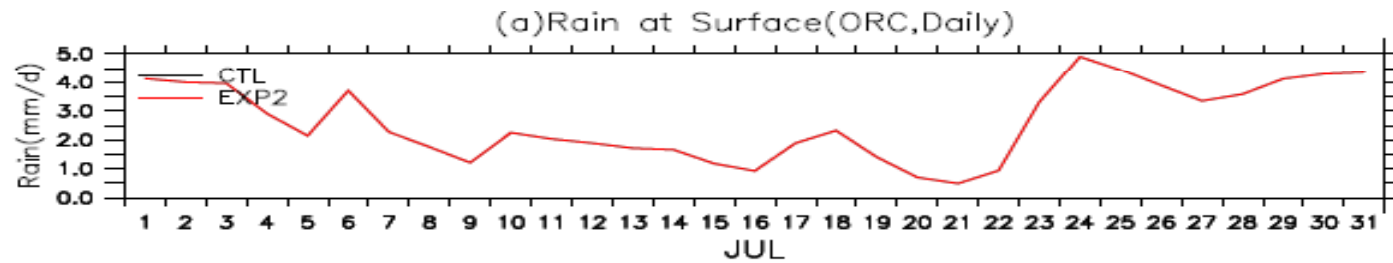
$$C_p(\theta, st) \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left[\underbrace{\lambda(\theta, st) \frac{\partial T}{\partial z}}_{\text{Conduction}} \right] - \underbrace{C_w \frac{\partial q_L T}{\partial z}}_{\text{Convection}} - S$$

The Effects of Liquid Water Transfer in Soil: ORCHIDEE-11 Forced Mode

	Soil Heat Convection	Soil Discretization	Thermal Property
CTL	No	Revised (8 meters)	Constant
EXP	Yes		



60W-110W,
25N-45N



Conclusions

➤ The ORC-II soil thermal model was modified to conserve energy: **I.** thermal property; **II.** same soil layer for moisture & temp.; **III.** soil heat convection-conduction.

I. The Sahara, western Australia are mostly affected by soil thermal properties (soil texture effects), with the surface temperature changes $\sim 1.5\text{K}$ (maximum) in JJA.

II. For most regions, the annual average surface temperature (soil moisture) changes $< \pm 0.1\text{K}$ ($\pm 10\%$) due to the different soil vertical discretization.

III. In daily or 3h time scale, the soil temp. varies $\sim 0.05\text{K}$ with rainfall $\sim 0-5\text{mm/d}$ (US, Summer). The variation decreases at larger time scale (monthly).

➤ Others:

- The changing of discharge/runoff ($\sim \pm 10\%$ humid regions) due to revision of soil discretization (?)

- The consideration of soil freezing parameterization (?)



Thank You! Questions & Comments ?

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